Low Level Wind Shear: Invisible Enemy To Pilots

On the afternoon of August 2, 1985, a landmark aircraft accident occurred at the Dallas/Fort Worth (DFW) airport. The tragic accident, which killed 137 of the 163 passengers on board Delta Airlines Flight 191, was responsible for making “wind shear” a more commonly known weather phenomenon and implementing many new changes with regard to wind shear detection (Ref. 1). On that day, thunderstorms were in the area of approach to runway 17L at the DFW International Airport, with a thunderstorm rain shaft right in the path of final approach. The crew decided to proceed through the thunderstorm, which turned out to be a critical error. Shortly after entering the storm, turbulence increased and the L1011 aircraft encountered a 26 knot headwind. Just as suddenly, the wind switched to a 46 knot tailwind, resulting in a loss of 72 knots of airspeed. This much of an airspeed loss on final approach, when the jet was only 800 feet above the surface, was unrecoverable and the aircraft eventually crashed short of the runway (Ref. 1). The sudden change in wind speed and direction that the aircraft encountered is called wind shear.

Wind shear can occur at many different levels of the atmosphere, however it is most dangerous at the low levels, as a sudden loss of airspeed and altitude can occur. Plenty of altitude is normally needed to recover from the stall produced by the abrupt change in wind speed and direction, which is why pilots need to be aware of the hazards and mitigation of low-level wind shear. So how does the phenomenon occur? There are a couple of ways, with the most dangerous coming from thunderstorm updrafts and downdrafts. The lifecycle of a thunderstorm starts with just an updraft as warm moist air at the surface rises in an unstable atmosphere to heights as high as 50 to 60 thousand feet in the air. As the updraft matures, precipitation forms and falls out as rain. This is when the downdraft portion of the thunderstorm occurs. Rain-cooled air sinks along with the downward drag from falling precipitation. The harder the rain falls, the stronger the downdraft can become. Also, stronger atmospheric winds aloft are sometimes able to be pulled down to the surface by the already established downdraft, adding to the intensity of the downward moving winds. The image below (Image 1) shows the updraft/downdraft interface of a thunderstorm where low-level wind shear might suddenly occur.
Another way that low-level wind shear can occur is when a “nocturnal” or nighttime inversion sets up as the lowest levels of the atmosphere “decouple” from the upper levels and stronger winds. A nocturnal inversion is basically a layer of warm air that develops in the lowest few thousand feet of the atmosphere as the planetary boundary layer becomes decoupled or “disconnected” from the stronger winds just above it. The layer of warm air marks the interface between the calmer and much cooler layer of air near the surface and the stronger winds and more well-mixed atmosphere above the inversion. This interface is another area where wind shear can occur. Many times, the change in temperature with height between the inversion and the atmosphere above it can create the pressure gradient necessary for a low level wind maximum between 25 (weaker) and 60 (stronger) knots. So, imagine taking off on a calm evening and climbing through 2 thousand feet. Suddenly, you encounter serious turbulence as a headwind or tailwind of 40 knots rocks your aircraft and changes your airspeed dramatically. This effect can quickly cause a pilot to lose control of the aircraft, especially if he or she is caught off guard or have never experienced the phenomenon before. To better understand the nocturnal inversion and low-level jet refer to the image below (Image 2).
The image above is a balloon sounding taken from Amarillo, Texas on the morning of September 1st, 2009. The squiggly white line where the left yellow oval is highlighted indicates temperature from the surface (bottom) up through the atmosphere. Notice the increase in temperature just above the surface, this is the nocturnal inversion. If you look at the wind barbs to the right (right yellow oval), you see an example of low-level wind shear. There is a south-southeasterly wind at the surface around 10 knots, then the wind increases to 35-40 knots out of the southwest over a short vertical distance. The temperature gradient caused by the nocturnal inversion helps to create the low-level jet around 40 knots in this case.

As mentioned above, many changes to how we detect and react to low-level wind shear have changed since the landmark crash of Flight 191. With the advent of the National Weather Service Doppler Radar, low-level velocity data has aided in the detection of low-level wind shear. Many large airports now have TDWRs (Terminal Doppler Weather Radar) installed near the airport to help detect changes in wind speed and direction. Also, at the time of the Flight 191 crash, the Federal Aviation Administration was in the process of testing and implementing LLWAS (Low Level Windshear Alert System). Although the new system had shown promise in detecting low-level wind shear in other incidents in recent
years, it was not able to detect wind shear before the Dallas/Fort Worth crash in 1985. As a result, the nation demanded something be done. In 1986, the FAA announced the National Integrated Windshear Plan. This included better training for pilots in how to detect and handle windshear situations and eventually led to the development of the TDWRs at airports today (Ref 1).

For more information on thunderstorm processes, low level wind shear, and forecasting techniques, including how the National Weather Service forecasts wind shear in Terminal Aerodrome Forecast’s go to: http://www.crh.noaa.gov/Image/lmk/aviation-conference/av-conf-talks/Wind_shear_final.pdf

References: